

還待進一步的診斷研究 (diagnostic study) 才能定論。

5. 致 謝

本研究是在國科會 NSC-64M-0202-1 (11) 計劃輔助下完成的。在計劃進行過程中，本系客座副教授陳泰然先生曾多次參與討論。有些觀念是在討論中獲得的。此外亦承蒙陳副教授幫忙分析觀測資料，此計劃才能如此順利完成。本系教授周根泉先生曾細心審閱初稿，提供不少寶貴意見。特此向他們表示由衷的感激。另外亦蒙中央氣象局及空軍氣象中心給予方便供給所需的資料在此並向有關人員致謝。

6. 參 考 資 料

- Bleck, R., 1965: Lineare approximations methoden zur bestimmung ein- und zweidimener filter des dynamischen meterlogie. Institut fur Theoretische, Meteorologie der Freien Universitate, Berlin.
- Cressman, G. P., 1959: An operational objective analysis system, *Mon. Wea. Rev.*, 87, 367-374.
- Haltiner, G. J., 1971: *Numerical weather prediction*, Wiley, New York.
- Holton, J. R., 1972: *An introduction to dynamic meteorology*, Academic Press, New York, 102-118.
- National Weather Service, 1972: Forecasters handbook No. 1: facsimile products, 9-26-9-31.
- Petterssen, S., 1956: *Weather analysis and forecasting*, McGraw-Hill, New York, 320-339.
- Saucier, W. J., 1955: *Principles of meteorological analysis*, University of Chicago Press, Chicago III., 29-38.
- Shuman, F. G., 1957: Numerical methods in weather prediction, II smoothing and filtering, *Mon. Wea. Rev.*, 85, 357-361.
- Spiegler, D. B., 1960: An analysis of barotropic forecast errors in cases of rapid sea level cyclogenesis, *Mon Wea. Rev.*, 88, 263-268.
- Thompson, P. D., 1961: *Numerical weather analysis and prediction*, Macmillan, New York.
- U. S. Air Force, 1966: Operational navigation chart, 1: 1,000,000, The Aeronautical Chart and Information Center, U. S. Air Force, St. Louis Mo.
- 胡仲英, 1974: 正壓大氣模式及其應用, 大氣科學 (中國氣象學會), 2, 1-7。
- 劉廣英及鄧施人, 1973: 利用電子計算機做數值天氣預報之研究 (下), 氣象預報與分析 (空軍氣象聯隊), 56, 6-8。
- 蔡清彥及胡仲英, 1975: 以相當正壓模式研究地形的動力效果。大氣科學, 2, 63-67。
- 蔡清彥, 1975: 數值天氣預報在臺灣地區的應用。國立臺灣大學大氣科學系, 數值天氣預報研究報告 NWP 01 號。

An Equivalent Barotropic Model and Its Forecast Errors in the Area of Taiwan and Its Vicinity

Ching-Yen Tsay

Dept. of Atmos. Sci., National Taiwan University, Taipei, Taiwan

Abstract

Numerical method for an equivalent barotropic model which includes grid structure, topography, finite difference equation, boundary conditions, data analysis and smoothing methods; are presented.

Three cases have been selected for the study of the limited-area

equivalent barotropic model. Case A predicts the movements of 500mb troughs and ridges related to cold-air outbreak in Taiwan. Case B also predicts the movements of troughs and ridges related to cold-air outbreak. The trough in this case moved after the cut-off low system opened up. Case C studies the movement of a trough in the eastern slope of Tibetan Plateau. Generally, movements of troughs and ridges in 24 hours were well predicted by the equivalent barotropic model. However, several systematic errors are apparently presented in the predictions after 36 hours. The most serious error is the unrealistic southward movement of a positive vorticity center during a trough passes over the northeastern part of Tibetan Plateau. As a result, an unrealistic low system is separated from the main trough and stays in Southeast Mainland. The error is suspected as the result of the model without considering the blocking, deflecting and frictional effect of Tibetan Plateau. Another systematic error is the displacement of forecast trough being fast in the north and slow in the south. The error may come from the beta-plane approximation in assuming the Coriolis parameter being constant. The geostrophic wind computed by this assumption is larger in the north and smaller in the south than its real value. The equivalent barotropic model did not predict the open up of cut-off low in case B but well predicted its movement after it opened up. The error is suspected partly due to the equivalent barotropic model being unable to predict the intensification, but mainly due to the scales of the initial low pressure centers being small and not able to be observed correctly by the model.

In the period of case C, a surface cyclone is formed and developed in Central East China. Synoptic analysis suggests that the formation of the cyclone is initiated by the positive vorticity advection ahead of 500mb trough, although the warm advection and moist instability in 1000—500mb layer may play important roles in the development of the cyclone.